ETHYLENE ACCUMULATION IN SOUTHEASTERN COASTAL PLAIN SOILS; SOIL CHARACTERISTICS AND OXIDATIVE-REDUCTIVE INVOLVEMENT

KEY WORDS: Soil-ethylene, organic matter, pH, nitrate content, soil matric potential

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## ABSTRACT

Two laboratory experiments were conducted to assess the extent of ethylene production in southeastern Coastal Plain soils; the influence of soil characteristics on ethylene accumulation; and the impact of oxidative, reductive, and oxidative-reductive interactive conditions upon ethylene formation. Soils from 21 sites over 9 selected soil series typical of the Coastal Plain physiographic region were air dried, sieved, incubated at 25°C with 50 mb matric potential or saturated soil water contents, and analyzed for ethylene accumulation. Soil nitrate, organic matter, and pH vs. ethylene accumulation were statistically evaluated by a quadratic surface response curve; a R<sup>2</sup> of 0.77 was obtained. In the second study, an oxidative-reductive interactive vat system was used to investigate the involvement of oxidative, reductive,

and oxidative-reductive conditions. Ethylene was found to accumulate under both oxidative and reductive conditions; however, the mixing of oxidative and reductive solutions did not increase the accumulation of ethylene.

### INTRODUCTION

Ethylene production in soils has been an area of considerable interest to researchers around the world for over ten years (Smith and Russel, 1969; Eplee, 1975; Hunt et al., 1980a). Ethylene is a plant hormone and is known to cause plant physiological responses such as epinastic growth of tomato at >2ppm concentration in the soil (Jackson and Campbell, 1975), and it has considerable potential importance in soil systems. Yet, the actual mode of  $^{\rm C}_2{\rm H}_4$  production in the soil and the conditions under which it is formed are not agreed upon. Nitrate has been reported to suppress  $^{\rm C}_2{\rm H}_4$  production (Smith et al., 1976; Hunt et al., 1980a). Goodlass and Smith, (1978), however, reported only very high concentration of nitrate affected  $^{\rm C}_2{\rm H}_4$  production in the soils they investigated. Hunt et al. (1980a) proposed that this difference could be related to physiochemical differences involving soil pH, Eh, Fe, and Mn.

Some authors believe  $\mathrm{C_2H_4}$  is produced under aerobic conditions (Lynch, 1975; Primrose, 1976), while others believe it to be produced anaerobically (Smith and Cook, 1974). Hunt et al. (1980a) proposed that  $\mathrm{C_2H_4}$  production in a Norfolk sandy loam soil may occur under both aerobic and anaerobic conditions. They found that aerobic accumulation of  $\mathrm{C_2H_4}$  was enhanced greatly by methionine, but anaerobic accumulation was affected very little by

methionine. The existence of the two processes would explain a considerable amount of disagreement in the literature concerning whether  ${\rm C_2H_4}$  production was an aerobic process or anaerobic process. However, an interactive oxidative-reductive process as suggested by Lynch (1975) cannot be excluded. The substrate for aerobic production of  ${\rm C_2H_4}$  could be produced under anaerobic conditions, but utilized for  ${\rm C_2H_4}$  production under aerobic conditions. A similar but reversed process is well documented for the interaction of oxidative and reductive conditions in nitrification and denitrification in wetland environments (Patrick and Mikkelsen, 1971).

This laboratory study was conducted to 1) assess the extent of  $^{\rm C}_2{}^{\rm H}_4$  production in southeastern Coastal Plain soils, 2) assess the influence of soil characteristics on  $^{\rm C}_2{}^{\rm H}_4$  accumulation, and 3) assess the impact of oxidative, reductive, and oxidative-reductive interactive conditions upon  $^{\rm C}_2{}^{\rm H}_4$  formation.

## METHODS AND MATERIALS

## Experiment I

A study was conducted with soil from 21 sites over 9 selected soil series typical of the Coastal Plain physiographic region. The soils were air dried and sieved through a 10 mesh sieve. They were analyzed for nitrate by copper sulphate extraction and specific ion electrode analysis (Lewis, 1961). The organic matter content was determined by weight loss ignition (Jackson, 1958), and pH was determined on a one to one soil to water suspension (W/V). Soil water content at 50 mb matric potential was deter-

mined on disturbed cores by the method described by Richards (1948). Saturation was considered to be the moisture content at which a sheen was left after pulling a spatula across the soil surface. Bulk density was determined by the core method. Properties of the soils are reported in Table 1.

Triplicate, twenty gram samples of each soil were placed in 37 ml glass bottles, sealed with rubber septum and incubated at 25°C. One set was incubated at 50 mb matric potential and one set was incubated under saturated water conditions. They were sampled by use of gas tight syringes for  $0_2$  and  $0_2$  and  $0_4$  on days  $0_4$ ,  $0_4$ ,  $0_4$ , and  $0_4$ ,  $0_4$ , and  $0_4$ ,  $0_4$ ,  $0_4$ ,  $0_4$ , and  $0_4$ ,  $0_4$ ,  $0_4$ ,  $0_4$ ,  $0_4$ , and  $0_4$ ,  $0_4$ ,  $0_4$ ,  $0_4$ , and  $0_4$ ,  $0_4$ ,  $0_4$ ,  $0_4$ ,  $0_4$ , and  $0_4$ ,  $0_4$ 

TABLE 1 Soil Characteristics and Maximum  ${\rm C_2H_{\Delta}}$  Accumulation

OBS	Soil	Organic Matter	pН	[NO <sub>3</sub> ]	[C2H4,max]
No.	Series	%		ppm	<sup>2</sup> ppm
1	Coxville	3.34	5.4	4.19	0.27
2 3	Dunbar	3.03	5.1	4.36	0.93
3	Duplin	2.44	5.6	1.32	1.31
4	Duplin	3.25	5.7	2.97	5.33
5	Goldsboro	2.11	5.4	2.23	0.93
6	Goldsboro	1.71	5.7	2.03	1.30
7	Portsmouth	1.70	5.5	1.76	1.15
8	Norfolk	1.53	5.8	2.46	1.22
9	Norfolk	1.07	6.1	3.27	0.27
10	Norfolk	1.38	6.1	1.32	2.07
11	Rains	3.10	5.5	2.70	2.84
12	Rains	2.05	5.7	1.60	2.10
13	Rains	2.27	5.7	4.79	0.83
14	Troup	0.70	5.9	1.09	1.02
15	Troup	0.81	6.2	1.66	2.06
16	Wagram	0.71	5.8	1.32	1.02
17	Wagram	1.19	6.1	1.32	0.17
18	Wagram	0.72	6.1	1.85	0.25
19	Wagram	1.21	6.3	3.78	0.93
20	Wagram	1.13	5.9	2.84	0.27
21	Wagram	0.69	5.8	1.39	2.27

saturation treatments, respectively. Ethylene concentrations were determined on a Hewlett Packard 5750 gas chromatograph equipped with a flame ionization detector and an aluminia F column at  $100\,^{\circ}\text{C}$ . Oxygen was determined by use of a Cd-Au electrode. The maximum  $^{\text{C}}_2\text{H}_4$  accumulation vs. soil nitrate, organic matter, and pH relationship was analyzed by a quadratic surface response curve.

## Experiment, II

An oxidative-reductive interactive reaction vat system as reported by Hunt et al., (1980b) was used. The system is a modification of the reaction vat system reported by Patrick et al., (1973) for studies of physicochemical processes under varying oxidative-reductive conditions. It allowed gas samples from oxidized and reduced soil suspensions to be obtained, and it allowed the liquid phase of the two systems to be interchanged. A Norfolk sandy loam (Typic paleudult) containing 75% sand, 22% silt, and 3% clay was passed through a 200 mesh sieve. Its pH was 5.2, and its organic matter content was 1.3%. Two hundred grams of the sieved soil and 1g of ground bermudagrass (Cynodon dactylon L.) were suspended in 2.5 liters of distilled water by a magnetic stirrer. The gas space above the suspensions of one vat contained air and the gas space above the other contained  ${\rm N}_2$ . These atmospheres were replaced daily to maintain distinct oxidized and reduced conditions. Duplicate oxidative-reductive systems were used. The gas spaces were sampled for  $0_2$  and  $C_2H_4$ . James solid state pH electrodes were used to measure pH. A calomel and a bright platinum electrode were used for measuring Eh. The liquid contents of the oxidative and reductive vats were interchanged

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after distinctly oxidized and reduced conditions had developed. The liquid phases were interchanged for about 10 min/24 h, which allowed for approximately 10% of the contents to be interchanged each day. In separate experiments, either 10 or 30 ppm of  $^{\rm C}_2{}^{\rm H}_4$  was added to the air space of the oxidized and reduced vats, and  $^{\rm C}_2{}^{\rm H}_4$  was monitored.

## RESULTS AND DISCUSSION

## Experiment I

The maximum  $C_2H_4$  accumulation for the soils incubated at 50 mb matric potential was less than the maximum  $C_2H_4$  accumulation in the same soils incubated at saturation (Fig. 1). This is in agreement with data of Hunt et al. (1980a), where significant  $C_2H_4$  accumulation only occurred at water contents very close to saturation. This indicates that  $C_2H_4$  accumulation potential in many soils of the southeastern Coastal Plain physiographic region is similar to that of many soils of the world as related to high soil water and low soil  $O_2$  concentrations. The average maximum accumulation was on day 7. However, it must be remembered that both production and consumption of  $C_2H_4$  are involved in accumulations. We have found accumulation increases followed by concentration declines to be common.

Data in Table 1 shows that these soils accumulated a maximum and a minimum of 5.33 and 0.17 ppm, respectively, under saturated conditions. Thus, all of the tested soils produced some  ${\rm C_2H_4}$ , and more than one-half of the soils tested accumulated more than 1.0 ppm. On part per million in these bottles equals 2.2 ng  ${\rm C_2H_4/g}$  of

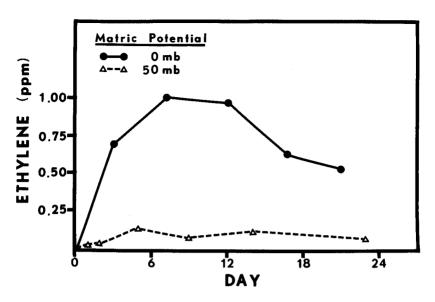


FIG. 1 Ethylene accumulation under variable matric potential.

soil. This level of  ${^{\rm C}2^{\rm H}}_4$  production could lead to damage of sensitive crops grown on these soils (Jackson et al., 1975, 1976, 1978).

Maximum  $^{\rm C}_2{}^{\rm H}_4$  production was correlated to soil organic matter, nitrate concentration and pH with a quadradic surface response; an  $^{\rm R}^2$  of 0.77 was obtained (equation 1).

$$C_{2}H_{4} = 487.40 - 71.41(0.M.) - 154.78 \text{ (pH)}$$
 [1]  
+ 14.75 (NO<sub>3</sub>) + 11.53 (0.M. x pH)  
- 0.84 (0.M. x NO<sub>3</sub>) - 2.28 (pH x NO<sub>3</sub>)  
+ 2.21 (0.M.)<sup>2</sup> + 12.27 (pH)<sup>2</sup> - 0.05 (NO<sub>3</sub>)<sup>2</sup>  
 $C_{2}H_{4} = \text{maximum ethylene, ppm}$   
0.M. = organic matter, % (0.5 to 3.5)  
NO<sub>3</sub> = nitrate concentration, ppm (1 to 5.4)  
pH = 5.0 to 6.3

The small  $R^2$  indicates that factors other than those tested were involved. Yet; one such factor, bulk density, which ranged from 1.04 to 1.85 did not improve the  $R^2$  of equation 1. Probably the nature of the organic matter (ease of degradation and chemical composition), iron, and manganese content, and initial microbial populations also affect  $C_2^{\rm H}_4$  accumulation. However, the quadratic surface response analysis gave much better correlation with our data than did simple linear regressions. This analysis suggests why there is disagreement in the literature concerning the response nature of various soils for  $C_2^{\rm H}_4$  accumulation when compared only by simple linear regression (Goodlass, 1978).

Although the data base is too limited to allow use of equation 1 as a broad-based predictor of soil  $C_2H_4$ , at lower organic matter contents (<1.04%),  $C_2H_4$  accumulation increased with lower pH values, and generally increased with increased NO $_3$  content in the studied range. At higher contents of organic matter (2.5 to 3.5%),  $C_2H_4$  was increased by increased pH and reduced NO $_3$ . The differential nature of  $C_2H_4$  accumulation with high and low organic matter could be interpreted to support the thesis of multiple  $C_2H_4$  production processes.

# Experiment II

The Eh of the reduced suspension (Fig. 2) decreased to less than 0 mV within the first 24 hour, and the reduced condition was maintained throughout the experiment. Conversely, the Eh of the oxidized suspension was continuously positive throughout the experiment (>+250 mV). The air space of the oxidized vat was

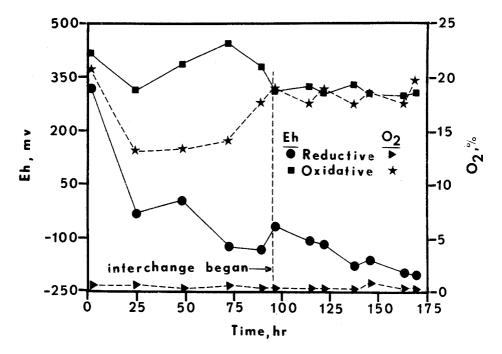


FIG. 2 Eh and  $\theta_2$  conditions in oxidative and reductive vats before and after solution interchange.

maintained above 12%  $0_2$ ; whereas, it was 0 to 1% in the reduced vat. The pH ranged from 6.4 to 7.0.

Figure 3 shows the accumulation of  ${\rm C_2H_4}$  for each vat over a 7-day period. The reduced vats were clearly accumulating more  ${\rm C_2H_4}$  than were the oxidized vats. This might be caused by a greater rate of  ${\rm C_2H_4}$  degradation under oxidative conditions as reported by Cornforth (1975). However, in a separate study when both 10 and 30 ppm of  ${\rm C_2H_4}$  were added to the vat head space, approximately 50% of the  ${\rm C_2H_4}$  was consumed under either oxidized or reduced condition. Similar results were found when the soil was incubated in small bottles with  ${\rm C_2H_4}$  added to the air space

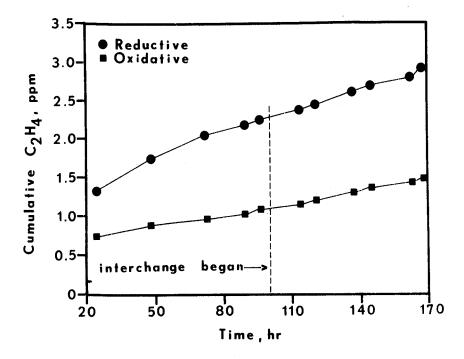


FIG. 3 Ethylene accumulation in oxidative and reductive vats befor and after solution interchange.

containing either air or  $N_2$ .

The lack of slope change in the  ${\rm C_2H_4}$  accumulation curve upon mixing of the oxidized and reduced liquid phases suggest that the accumulation of  ${\rm C_2H_4}$  is not dependent upon or stimulated by the interaction of substances accumulated under oxidized or reduced conditions. Since our mixing method was a batch exchange, we cannot rule out the involvement of unstable intermediates at the oxidative-reductive interfaces of the soil. However, our data suggest that soil accumulation of  ${\rm C_2H_4}$  can occur in the Norfolk loamy sand in either oxidized or reduced microsite without interchange.

From these laboratory experiments the authors concluded: 1)  $^{\rm C}_2{}^{\rm H}_4$  is commonly produced by soils of the southeastern Coastal Plain. 2) The extent of  $^{\rm C}_2{}^{\rm H}_4$  accumulation was correlated with an  $^{\rm R}^2$  of 0.77 to soil organic matter, pH, and nitrate concentration. 3)  $^{\rm C}_2{}^{\rm H}_4$  can accumulate under both oxidative and reductive conditions, and it is not greatly stimulated by batch interchange of oxidized or reduced suspension of Norfolk loamy sand. Its occurrence in both oxidized and reduced conditions would support the idea of two independent processes for  $^{\rm C}_2{}^{\rm H}_4$  production and accumulation in soils. The existence of at least two processes would explain conflicts in the literature concerning whether  $^{\rm C}_2{}^{\rm H}_4$  is formed aerobically or anaerobically in soils.

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